



Development of the GEOS-5 atmospheric general circulation model: evolution from MERRA to MERRA2

A. Molod¹, L. Takacs², M. Suarez³, and J. Bacmeister⁴

¹University of Maryland College Park, College Park, MD, USA

²Science Systems and Applications, Inc., Lanham, MD, USA

³NASA Goddard Space Flight Center, Greenbelt, MD, USA

⁴National Center for Atmospheric Research, Boulder, CO, USA

Correspondence to: A. Molod (andrea.molod@nasa.gov)

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Abstract. The Modern-Era Retrospective Analysis for Research and Applications-2 (MERRA2) version of the Goddard Earth Observing System-5 (GEOS-5) atmospheric general circulation model (AGCM) is currently in use in the NASA Global Modeling and Assimilation Office (GMAO) at a wide range of resolutions for a variety of applications. Details of the changes in parameterizations subsequent to the version in the original MERRA reanalysis are presented here. Results of a series of atmosphere-only sensitivity studies are shown to demonstrate changes in simulated climate associated with specific changes in physical parameterizations, and the impact of the newly implemented resolution-aware behavior on simulations at different resolutions is demonstrated. The GEOS-5 AGCM presented here is the model used as part of the GMAO MERRA2 reanalysis, global mesoscale simulations at 10 km resolution through 1.5 km resolution, the real-time numerical weather prediction system, and for atmosphere-only, coupled ocean-atmosphere and coupled atmosphere-chemistry simulations.

The seasonal mean climate of the MERRA2 version of the GEOS-5 AGCM represents a substantial improvement over the simulated climate of the MERRA version at all resolutions and for all applications. Fundamental improvements in simulated climate are associated with the increased re-evaporation of frozen precipitation and cloud condensate, resulting in a wetter atmosphere. Improvements in simulated climate are also shown to be attributable to changes in the background gravity wave drag, and to upgrades in the relationship between the ocean surface stress and the ocean roughness. The series of resolution-aware parameters related

to the moist physics was shown to result in improvements at higher resolutions and result in AGCM simulations that exhibit seamless behavior across different resolutions and applications.

1 Introduction

The various activities of NASA's Global Modeling and Assimilation Office (GMAO) necessitate a model that can function seamlessly across many different resolutions and applications. These applications include real-time atmospheric analyses and forecasts at a resolution of 0.25°, long-term reanalyses at 0.5°, coupled atmosphere–ocean and coupled atmosphere–chemistry simulations at 1 or 2°, and global mesoscale simulations at 7 km and higher. The Modern-Era Retrospective Analysis for Research and Applications-2 (MERRA2) version of the Goddard Earth Observing System-5 (GEOS-5) atmospheric general circulation model (AGCM) is part of an ongoing development of a new generation AGCM at GMAO. The focus of the development of the MERRA version of the GEOS-5 AGCM was on the behavior of the components of the hydrological cycle in reanalysis mode, while the focus of the development of the MERRA2 AGCM was on a model that functions seamlessly in numerical weather prediction, reanalysis, climate, and global mesoscale modes. To this end, some of the physical parameterizations were replaced, some parameters governing the behavior of other physical parameterizations were changed, and

Table 1. Changes in GCM Algorithms from MERRA to MERRA2.

Module	Algorithm change	Comments
Moist	Increased re-evaporation of precipitation	Fundamental change in model climate
	Modified autoconversion	Fundamental change in model climate
	Modified effective radius of cloud drops	Fundamental change cloud forcing
	Anvil fractions cut in half	Fundamental change cloud forcing
	AutoConvert “warm fog”	Important change in coupled simulations
	New critical RH with resolution dependence	Substantial change in simulated moisture
	Cloud base set at planetary boundary layer (PBL) depth	Remove clouds detraining below PBL height
	RAS timescale no longer depends on turbulence	
	Stochastic RAS with resolution dependence	Substantial impact at high resolution
Turb	Remove restrictions on diffusion from Louis	Increase near surface diffusion
	Reformulate turbulent length scale in Louis	
	Reduce Lock scheme when there is wind shear	Impact on marine PBL
	Reduce cloud top entrainment for Lock plumes	Impact on marine PBL
Surf	Implement Helfand and Schubert scheme	Improve land temperatures
	Remove viscous sublayer over land surfaces	Reduce wind bias in S. Ocean
	Change ocean roughness for middle wind regimes	Increase tropical cyclone intensity
	Change ocean roughness high wind regimes	
Land surf	Changed parameters for evapotranspiration	Impact on ratio of surface to canopy evaporation
GW (Gravity Wave) drag	Changed profile of background drag	Substantial impact on QBO
	Added intermittency of drag	Impact on timing of winter jet breakup

resolution-aware parameters were implemented in the moist process parameterizations.

Many studies exist that describe major improvements in new versions of AGCMs and show the improvements in simulations as compared to reanalyses and other observations (i.e., Neale et al., 2013; Donner, et al., 2011; Pope et al., 2000). The present study adds to that type of analysis by carefully documenting the connection between individual changes in the physical parameterizations of the AGCM and improvements in the climate simulation at coarse resolution. A series of sensitivity experiments were conceived and analyzed to explore, step by step, each important change in parameterizations between the MERRA and MERRA2 AGCMs, and to demonstrate the impact on the simulated climate. The present study also describes and analyzes the improvements in high-resolution simulations due to some changes in parameterizations specifically targeted for those resolutions.

The details of the changes in the AGCM physical parameterizations are described in the next section, the step by step experiments to isolate the impacts of these changes are described in Sect. 3, and the impacts of the resolution-aware aspects of the AGCM are described in Sect. 4. The study is summarized in Sect. 5.

2 Description of the MERRA2 version of the GEOS-5 AGCM

The generation of the GEOS-5 AGCM that was used as part of NASA’s Modern-Era Retrospective Analysis for Research and Applications (MERRA) is described in Rienecker et al. (2008), and most of the subsequent development of the physical parameterizations for the current MERRA2 version is described in Molod et al. (2012). In addition to the changes in the physical parameterizations, the development of the MERRA2 AGCM also included two fundamental elements that will not be addressed in the present study. The horizontal discretization of the MERRA2 AGCM is computed on the cubed sphere grid of Putman and Lin (2007), although it still retains the option to use the latitude–longitude discretization. The cubed sphere grid allows for the relative uniformity of grid spacing at all latitudes, and avoids the grid spacing singularities found in the latitude–longitude grid. In addition, the MERRA2 AGCM has been modified to account for the change in total mass due to the change in total water content computed in the moist and turbulence processes. The total mass of each layer is adjusted to include these changes in total water content, and the associated adjustment is then made to the specific masses of all constituents, including water substances. The benefit for the AGCM mean simulated climate is small, but results in the conservation of dry mass during the simulation. The algorithm for this adjustment and the benefits for AGCM simulations and data assimilation experiments are described in detail in Takacs et al. (2014).

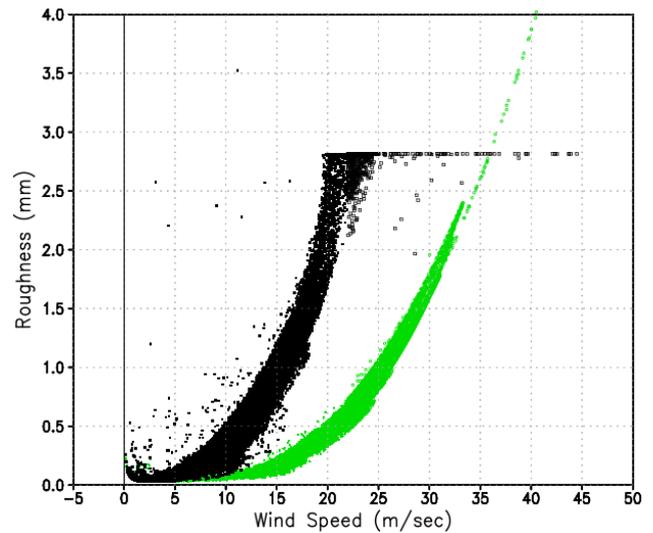
Table 2. Experiments to attribute MERRA to MERRA2 AGCM simulation changes to changes in parameterizations.

Experiment	Description
Control	
Exp. 1	Back off change in ocean roughness
Exp. 2	Exp. 1 + Back off gravity wave background drag and surface hydrology
Exp. 3	Exp. 2 + use old surface layer parameterization
Exp. 4	Exp. 3 + Back off increase of Richardson-number diffusion
Exp. 5	Exp. 4 + Back off decrease of critical RH aloft and decrease below
Exp. 6	Exp. 5 + Back off increase of all re-evaporation
Exp. 7	Exp. 6 + Back off gravity wave drag intermittency

A brief summary of the model's physical parameterizations relevant to the present study is provided here. The GEOS-5 AGCM physics includes parameterization schemes for atmospheric convection, large-scale precipitation and cloud cover, longwave and shortwave radiation, turbulence, gravity wave drag, a land surface model, a thermodynamic sea ice model, and a simple glacier model.

Convection is parameterized using the Relaxed Arakawa–Schubert (RAS) scheme of Moorthi and Suarez (1992) and includes a scheme for the generation and re-evaporation of falling rain (Bacmeister et al., 2006). A stochastic Tokioka trigger function (Bacmeister and Stephens, 2011) governs the lower limits on the allowable entrainment by sampling from a probability distribution function (PDF) with specified parameters. The prognostic cloud cover and cloud water and ice scheme is from Bacmeister et al. (2006). The scheme includes large-scale condensation governed by the PDF described in Molod (2012), evaporation, autoconversion and accretion of cloud water and ice, sedimentation of cloud ice, and re-evaporation of falling precipitation.

The turbulence parameterization that acts above the surface layer is based on the non-local scheme of Lock et al. (2000), acting together with the Richardson-number-based scheme of Louis and Geleyn (1982). The original Lock scheme was extended in GEOS-5 to include moist heating and entrainment in the unstable surface parcel calculations. The parameterization of surface layer turbulence is the Monin–Obukhov similarity theory based scheme described in Helfand and Schubert (1995), and includes the effects of a viscous sublayer for heat and moisture transport over all surfaces except land. The ocean roughness is determined by a polynomial which is a blend of the algorithms of Large and Pond (1981) and Kondo (1975), modified in the mid-range wind regime based on recent observations in the Southern

**Figure 1.** Scatter diagram of surface wind speed (m s^{-1}) vs. ocean roughness (mm) in the MERRA (green) and MERRA2 (black) AGCMs.

Ocean according to Garfinkel et al. (2011) and in the high wind regime according to Molod et al. (2013).

The longwave radiative processes are described by Chou and Suarez (1994), and the shortwave radiative processes are from Chou and Suarez (1999). The gravity wave parameterization computes the momentum and heat deposition into the grid-scale flow due to orographic (McFarlane, 1987) and non-orographic (after Garcia and Boville, 1994) gravity wave breaking. The land surface model from Koster et al. (2000) is a catchment-based scheme that treats subgrid-scale heterogeneity in surface moisture statistically. Glacial thermodynamic process are parameterized using an adaptation of the Stieglitz et al. (2001) snow model to glacial ice (Culather et al., 2014), and the catchment and glacier models are each coupled to the multi-layer snow model of Stieglitz et al. (2001). Sea ice albedos in the Northern Hemisphere are from the monthly mean observations of Duyenkerke and de Roode (2001).

3 Evolution of low-resolution simulated climate from MERRA AGCM to MERRA2 AGCM

The mean climate characteristics of a single 30-year MERRA2 AGCM simulation on the latitude–longitude grid at a horizontal resolution of $0.5^\circ \times 0.5^\circ$ were evaluated by comparison with reanalysis and with different satellite and in situ based observational estimates (Molod et al., 2012). They found substantial improvements in some key aspects of the mean circulation in the MERRA2 version of the GEOS-5 AGCM, and also reported on existing discrepancies between the modeled and observed climates. Here we present the results of a series of experiments designed to attribute each

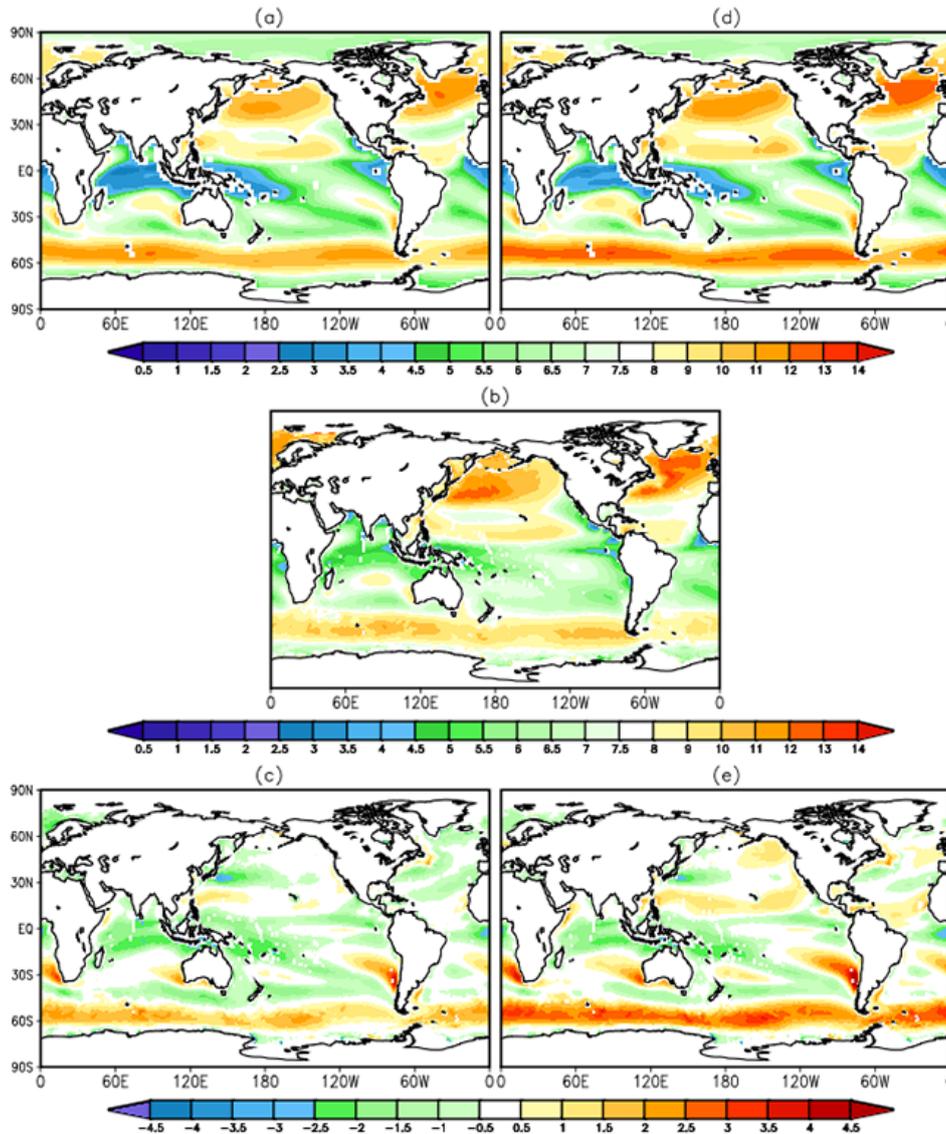


Figure 2. A 30-year average December-January-February surface wind speed (m s^{-1}) from (a) MERRA2 control, (b) GSSTF, (c) control-GSSTF, (d) experiment 1, (e) experiment 1-GSSTF.

fundamental improvement in AGCM simulated climate to a specific change in parameterization.

The experiments to be described in this section were all conducted on the latitude–longitude grid at $2.0 \times 2.5^\circ$ horizontal resolution, on a vertical hybrid eta–pressure coordinate grid with 72 levels, spaced to increase the resolution near the surface and near the tropopause, with an upper lid at 0.01 mb. The simulations were all forced with observed sea surface temperatures (Reynolds et al., 2002), and ran for 30 years each. The sequence of experiments was designed to start with the MERRA2 AGCM as the control and backtrack, one parameterization change at a time or small groups of parameterization changes as a time, to a model that replicates the MERRA AGCM simulated climate. The parameteriza-

tion changes are listed in Table 1, and the full sequence of the control and 7 sensitivity experiments to be described in this section is listed in Table 2.

3.1 Ocean surface winds

The parameterization of the surface layer turbulence in the MERRA2 AGCM includes a substantial modification of the functional relationship between ocean surface roughness and wind stress, shown in Fig. 1. The relationship for the moderate range surface wind speeds in the MERRA2 AGCM (black) is based on the implementation of Garfinkel et al. (2011), and the increase in roughness relative to the MERRA AGCM (green) is apparent in the figure. The first experiment in the series, experiment 1, reverts back to the formulation for

the relationship between ocean surface roughness and stress used in the MERRA AGCM. The effect for simulations at $2^\circ \times 2.5^\circ$ resolution is expected to be an increase in experiment 1's simulated surface wind speeds in the mid-range of wind speeds, that is, in the $5\text{--}25\text{ m s}^{-1}$ range. The relationship between wind speed and roughness for the higher wind speeds in the MERRA2 AGCM is based on Molod et al. (2013), and the figure shows reduction of roughness at speeds greater than approximately 30 m s^{-1} . This change is expected to result in a net increase of wind speeds in higher wind regimes, as was shown in Molod et al. (2013), for simulations at 0.25° resolution, but this impact is not apparent at the resolution of the experiments described here because the simulated wind speeds generally do not reach 30 m s^{-1} .

Surface wind speeds from the MERRA2 AGCM control and experiment 1 are shown in Fig. 2. The change in the simulated surface winds is most apparent in the Southern Hemisphere, where the Goddard Satellite-based Surface Turbulent Fluxes (GSSTF; Shie et al., 2009) surface winds, seen in Fig. 2b, show values near 8 m s^{-1} , experiment 1 (Fig. 2d) shows surface winds near 12 m s^{-1} , and the MERRA2 AGCM control (Fig. 2a) shows surface winds near 10 m s^{-1} . The difference from the GSSTF estimate shows a reduction in the difference from up to 4 m s^{-1} in experiment 1 (Fig. 2e) to a difference of up to 2 m s^{-1} in the MERRA2 AGCM control (Fig. 2c), pointing out the improvement in AGCM simulated climate due to the change in roughness formulation.

3.2 Quasi-biennial oscillation

The non-orographic gravity wave source in the MERRA (and MERRA2) AGCMs is specified as a function of latitude, and the source level is 400 mb, after Garcia and Boville (1994). The gravity wave source spectrum is specified a priori, and includes wave components with phase velocities in the range of -40 to 40 m s^{-1} , at intervals of 10 m s^{-1} . The latitudinal profile of background non-orographic drag in the MERRA2 AGCM was modified to include a source related to tropical precipitation in addition to the local maxima related to storm track precipitation. The use of this structure is based on the physical argument that the major sources for non-orographic gravity waves are convective and frontal systems. The MERRA2 AGCM profile was designed to mimic the behavior of more complex models of non-orographic gravity wave sources such as the state-dependant model of Richter et al. (2010), or the spectral wave source models of Manzini et al. (1997) and Scaife et al. (2002).

The latitudinal dependence of the background non-orographic drag profiles used in the MERRA2 and the MERRA AGCMs are shown in Fig. 3. Experiment 2 of the series examines the impact of the change in the gravity wave drag parameterization of background drag in the tropics and returns to the MERRA AGCM background drag. Experiment 2 is therefore expected to exhibit a stratospheric wind with

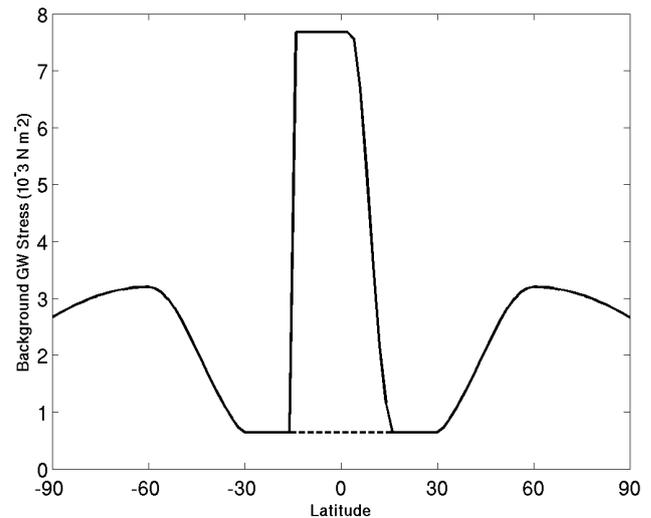


Figure 3. Background non-orographic drag from the MERRA (dashed line) and MERRA2 AGCM (solid line) simulations. The dashed line underlies the solid line outside of the tropics.

no quasi-biennial oscillation (QBO) variability. The tropical (latitude range 10° S to 10° N) zonal average of the zonal wind as a function of height and time is shown in Fig. 4. The patterns of large easterly and westerly winds that slant downwards in pressure as time proceeds indicate the downward propagation of the variations due to the QBO. The QBO pattern is seen in Fig. 4a and c, which are the results from experiment 1 from and MERRA reanalysis (Rienecker et al., 2011), respectively. Figure 4b, which shows the results of experiment 2 that uses the old background drag formulation (from Fig. 3), shows no QBO pattern of variability.

3.3 Stable surface layer fluxes

The parameterization of surface layer turbulence in the MERRA2 AGCM uses the scheme of Helfand and Schubert (1995) based on Monin–Obukhov (MO) similarity theory. The scheme replaced the Louis (1979) scheme that was used in the surface layer only (the bottom model level) in the MERRA AGCM. The fundamental differences between the Louis (1979) description of surface layer turbulence and the scheme of Helfand and Schubert (1995) are the stable surface layer stability functions and the formulation for the viscous sublayer (the laminar layer that can act to impede the flux of heat and moisture). The viscous sublayer in the Louis scheme is present over unvegetated land surfaces and over the oceans. The Helfand scheme removes the viscous sublayer over all land surfaces, including bare soil. The stable surface layer stability functions in the Helfand scheme are such that there is an increased turbulent heat exchange (of both signs) under stable conditions. Figure 5 shows a scatter diagram of the sensible heat flux as a function of the surface bulk Richardson number under conditions where the monthly

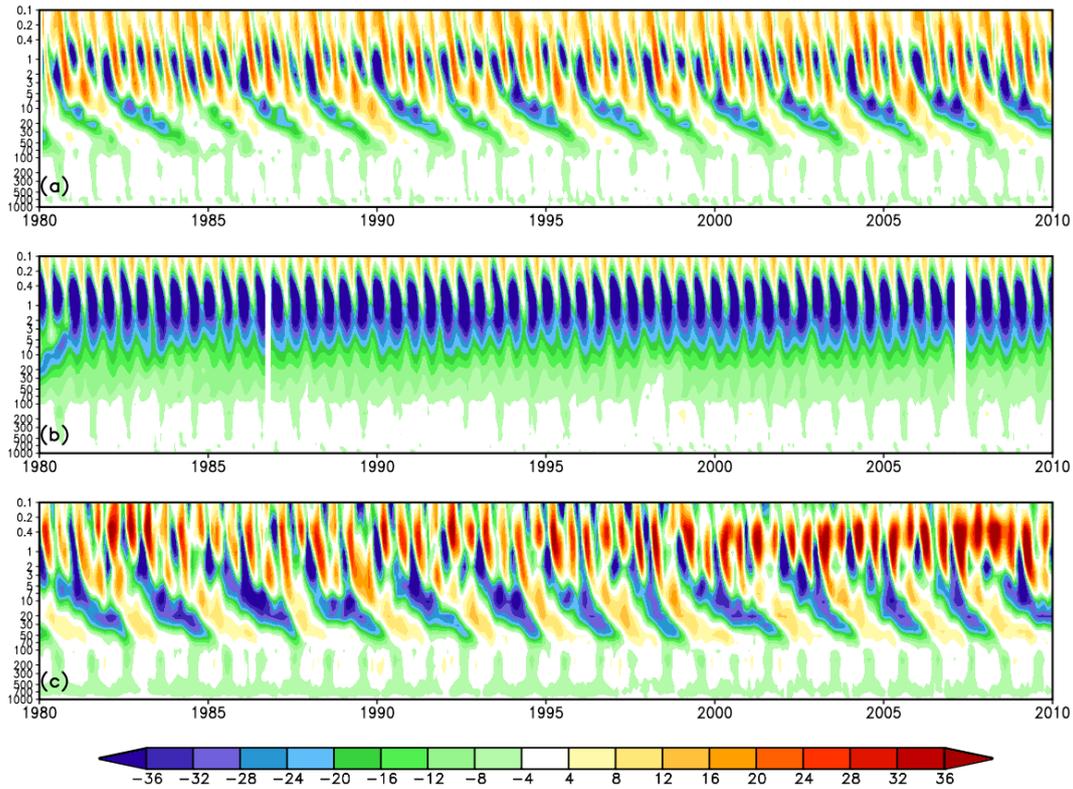


Figure 4. Spatial average of zonal wind in m s^{-1} from 10°S to 10°N latitude as a function of pressure level in millibars and time from (a) MERRA2 AGCM control, (b) experiment 2 and (c) MERRA reanalysis.

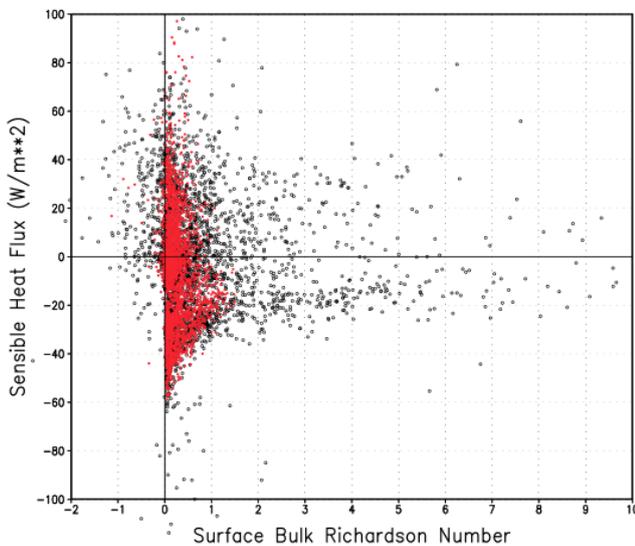


Figure 5. Surface bulk Richardson number as a function of sensible heat flux in W m^{-2} in a single July from experiment 3 (red) and experiment 2 (black).

mean air temperature exceeds the monthly mean skin temperature. The black points are from the simulation with the Helfand surface layer, and the red are from the simulation

using the surface layer from the Louis scheme. The larger values of sensible heat flux in the Helfand simulation are apparent, and even more apparent when the monthly mean sensible heat flux is downward.

Experiment 3 was designed to examine the impact of the change in the surface layer parameterization by reverting back from the Helfand and Schubert scheme to the Louis scheme. The turbulence parameterization above the surface layer, which is the combination of Lock et al. (2000) and Louis and Geleyn (1982) is the same in experiments 2 and 3. Figure 6 shows the sensible heat flux from experiments 2 and 3 along with the difference between them. The differences shown here are attributable to the removal of the viscous sublayer over land in the Helfand and Schubert scheme and to the change in the stable layer stability functions. Figure 6c shows that over most land surfaces the difference in sensible heat flux is negative, indicating less sensible heat flux when using the Louis scheme. This difference is consistent with the expectation that the removal of the viscous sublayer over land surfaces in the Helfand scheme lessens the resistance to turbulent exchange. Figure 6c also shows regions where the sensible heat flux is greater in the Louis scheme than in the Helfand scheme. These are regions where the surface layer is stable, and where the sensible heat itself is largely downward (i.e., the air temperature is greater than the

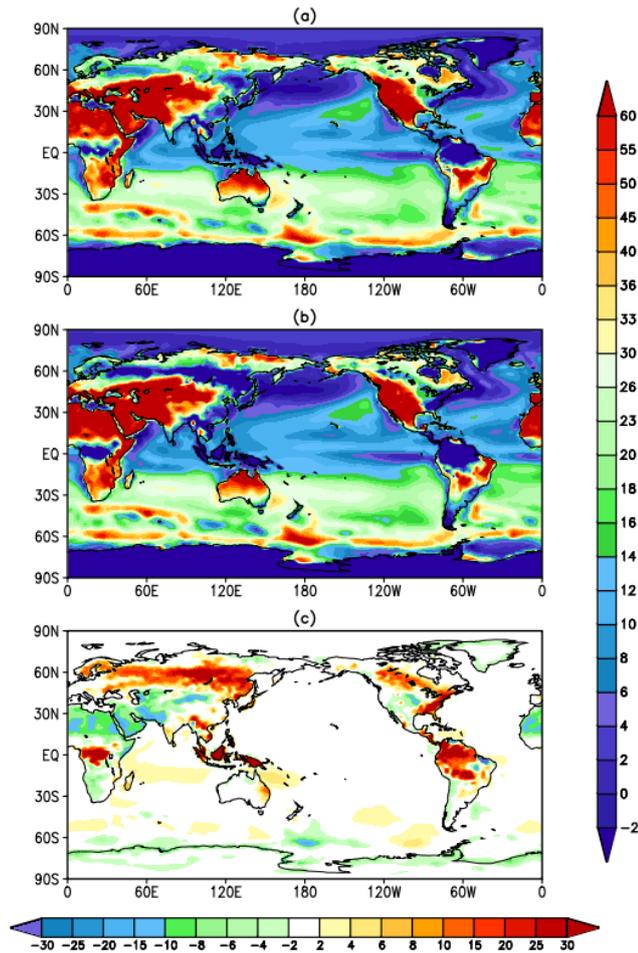


Figure 6. June-July-August-averaged sensible heat flux in W m^{-2} from (a) experiment 3 (Louis surface layer scheme), (b) experiment 2 (Helfand scheme) and (c) the difference of experiment 3 minus experiment 2.

skin temperature). The change in stability functions between the Louis and Helfand schemes, which allows more turbulent exchange in the Helfand scheme, is consistent with the sign of the difference in sensible heat flux in regions where the heat flux is downward.

3.4 Atmospheric moisture, clouds and stationary wave pattern

3.4.1 Critical relative humidity

The algorithm for large-scale condensation, as described in Bacmeister et al. (2006), assumes that the PDF of total water is top-hat shaped. The width of the PDF can be shown to be associated with a critical relative humidity (RH_{crit}) that governs cloud macrophysical and microphysical processes such as condensation and evaporation (Molod, 2012). The relationship between RH_{crit} and PDF width is such that a wider PDF corresponds to a lower RH_{crit} . The MERRA2

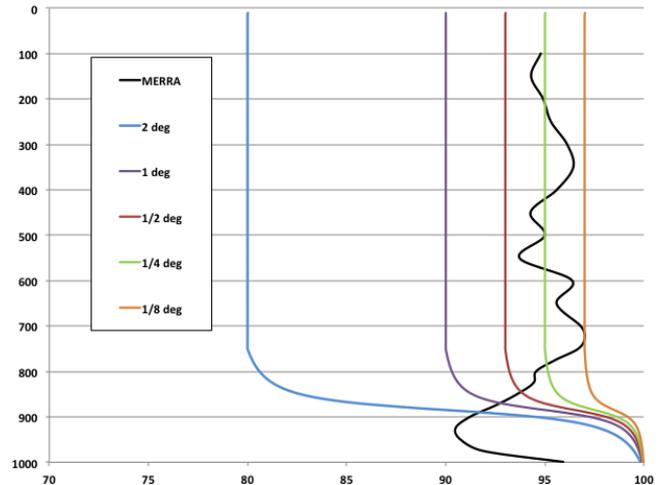


Figure 7. Critical relative humidity. Black from MERRA AGCM formulation, blue from MERRA2 AGCM formulation for 2-degree resolution, purple from MERRA2 AGCM formulation for 1-degree resolution, brown from MERRA2 AGCM formulation for 1/2-degree resolution, green from MERRA2 AGCM formulation for 1/4-degree resolution, orange from MERRA2 AGCM formulation for 1/8-degree resolution.

AGCM RH_{crit} (Molod, 2012) represents a change in both the magnitude and vertical structure from the RH_{crit} in the MERRA AGCM. Typical RH_{crit} profiles from the MERRA and MERRA2 AGCMs are shown in Fig. 7, and indicate generally lower values in the MERRA2 AGCM formulation except in the boundary layer, where turbulent mixing is sufficient to homogenize the total water distribution and so result in narrower PDF.

Experiment 5 was designed to examine the impact of the change in RH_{crit} . Removing this change, which for much of the atmosphere means a larger RH_{crit} , should result in a simulation that is generally wetter because the atmosphere is being adjusted back to a higher relative humidity (RH). The zonal mean RH from experiments 4 and 5, along with the difference between them, is shown in Fig. 8. The experiment 5 minus experiment 4 difference shows a clear increase in RH in the MERRA AGCM-like experiment due to the increase in RH_{crit} . Relative to available observational verification, the MERRA2 AGCM shows a general wet bias (Molod et al., 2012), which means that the RH field of experiment 4 is closer to the observational verification than the RH field of experiment 5.

In addition to having a substantial impact on atmospheric moisture, the change in RH_{crit} also had an impact on the distribution of cloud cover. The higher RH_{crit} above the boundary layer of the MERRA AGCM-like experiment may be expected to result in two potentially competing effects. The requirement that the atmosphere must have a higher RH before condensation can take place could result in reduced cloud cover. The higher RH_{crit} , however, also leaves behind an at-

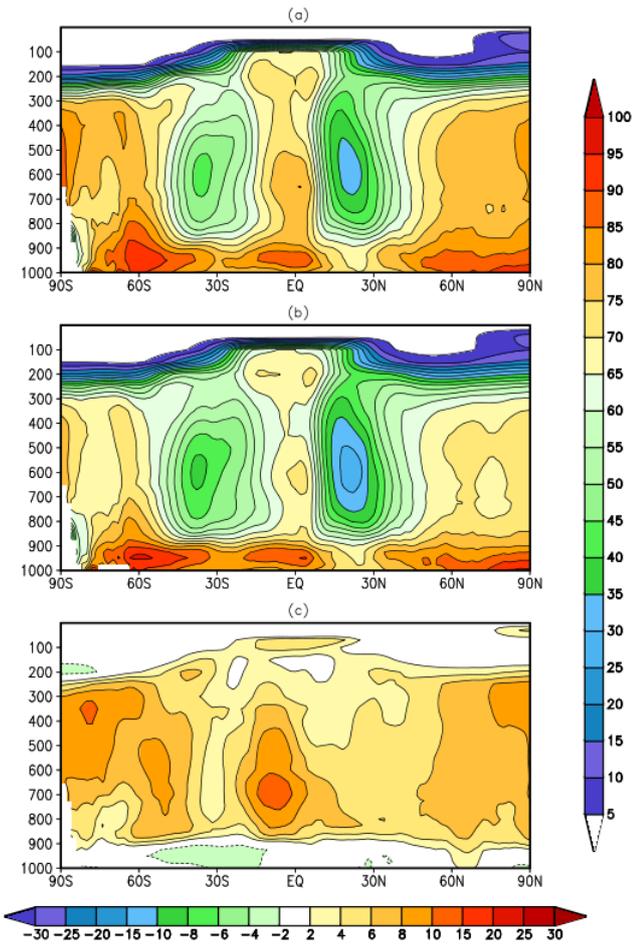


Figure 8. December-January-February-averaged relative humidity in percent from (a) experiment 5 (MERRA AGCM-like), (b) experiment 4 (MERRA2 AGCM-like) and (c) the difference of experiment 5 minus experiment 4.

mosphere with a higher RH (as seen in Fig. 8), which would be associated with increased cloud cover. One of these two effects may dominate, or they may cancel and the result would be no change in cloud cover.

Figure 9 shows the zonal mean cloud cover from experiments 4 and 5, and an observational estimate of zonal mean cloud cover from AIRS (Atmospheric Infrared Sounder). The MERRA AGCM-like experiment, experiment 5 (Fig. 9a), shows increased cloud cover in the 300–600 mb range relative to the MERRA2 AGCM-like experiment, in particular at high latitudes in both hemispheres. In this regard, the MERRA2 AGCM-like experiment result more closely resembles the AIRS cloud cover estimate (Fig. 9c).

The results shown here demonstrate that influence of the wetter atmosphere due to the higher RH_{crit} is the dominant of the two competing effects in the GEOS-5 AGCM. It is not clear, however, precisely why the atmospheric RH increases beyond the change in RH_{crit} . It may be speculated that it is related to the sequencing of the macro- and micro-physical

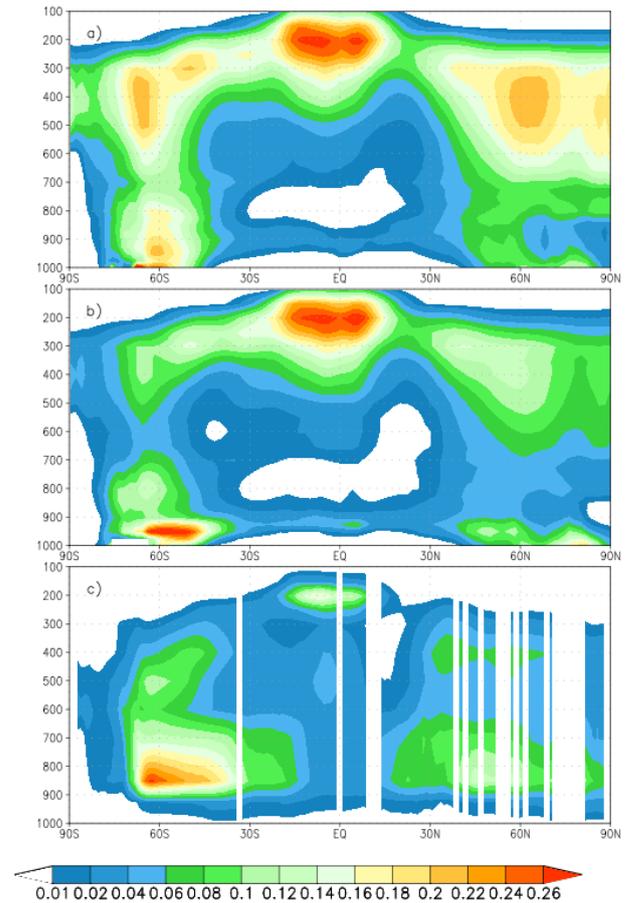


Figure 9. December-January-February-averaged cloud fraction from (a) experiment 5 (MERRA AGCM-like), (b) experiment 4 (MERRA2 AGCM-like) and (c) AIRS retrievals.

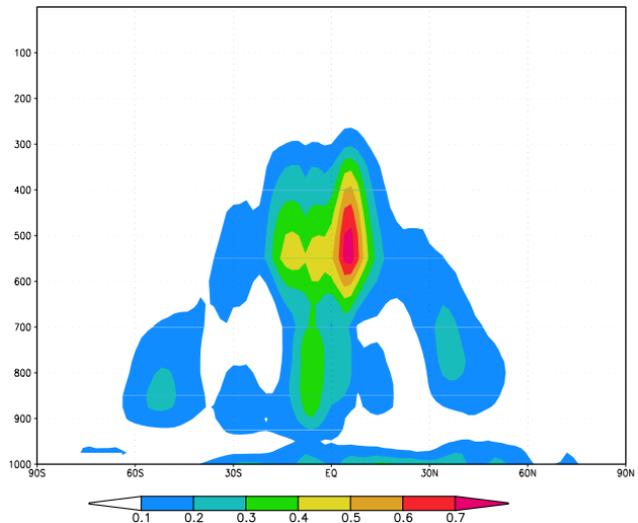


Figure 10. The difference (MERRA2 AGCM-like minus MERRA AGCM-like) of zonal mean specific humidity source term due to all re-evaporation for December-January-February.

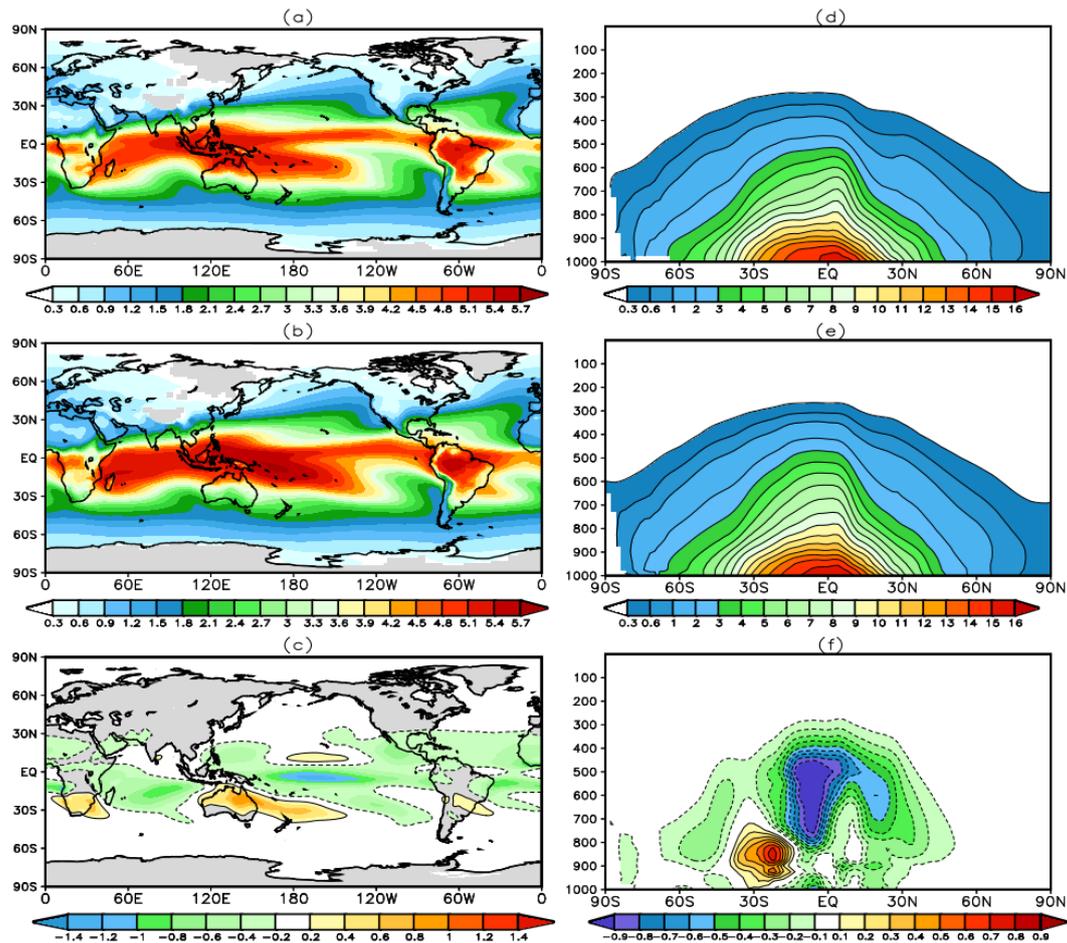


Figure 11. (a) December-January-February total precipitable water in millimeters from experiment 6 (MERRA AGCM-like), (b) same as (a) but from experiment 5, (c) same as (a) but the difference of experiment 6 minus experiment 5, (d) December-January-February-specific humidity in g kg^{-1} from experiment 6 (MERRA AGCM-like), (e) same as (d) but from experiment 5, (f) same as (d) but the difference of experiment 6 minus experiment 5.

processes in the moist parameterization combined with the enhanced rates of rain and condensate re-evaporation (see Sect. 3.4.2), leaving behind an atmosphere which is no longer in equilibrium with the specific RH_{crit} (wetter in our case). These sequencing issues are typical of AGCM moist parameterizations, as can be seen, for example, in the complex sequencing described in Gettelman et al. (2015).

The MERRA-like AGCM result (experiment 5) also shows smaller cloud cover near the boundary layer at almost all latitudes. At high latitudes, the MERRA2 AGCM-like boundary layer cloud more closely resembles the AIRS estimate, in the tropics the MERRA2 AGCM-like boundary layer cloud is larger than AIRS, while the MERRA AGCM-like boundary layer cloud is smaller. The change in boundary layer cloud between the MERRA AGCM-like experiment and the MERRA2 AGCM-like experiment is not consistent with the free atmosphere response to the RH_{crit} change because the presence of boundary layer turbulence makes RH_{crit} less of a determining factor for model mean relative humidity there.

3.4.2 Re-evaporation of precipitation and condensate

The MERRA2 AGCM scheme for the re-evaporation of precipitation and suspended cloud water and ice contains a series of new parameter settings that result in a substantial increase over the MERRA model in the re-evaporation of snow and ice. The impact of the changes in parameter settings on the water vapor source due to re-evaporation for the December-January-February average is shown in Fig. 10. The largest increase in the MERRA2 AGCM is aloft, near 500 mb, where the increase is up to $0.7 \text{ g kg}^{-1} \text{ day}^{-1}$.

Experiment 6 examines the impact of the change in re-evaporation of snow and ice in the MERRA2 AGCM, which is perhaps the most crucial parameterization change. The removal of this change is expected to result in a drier atmosphere, in particular aloft. Because of the importance of this change in parameterization, this simulation is expected to resemble in large part the climatology of the AGCM used as part of MERRA. Figure 11 shows the direct impact of the

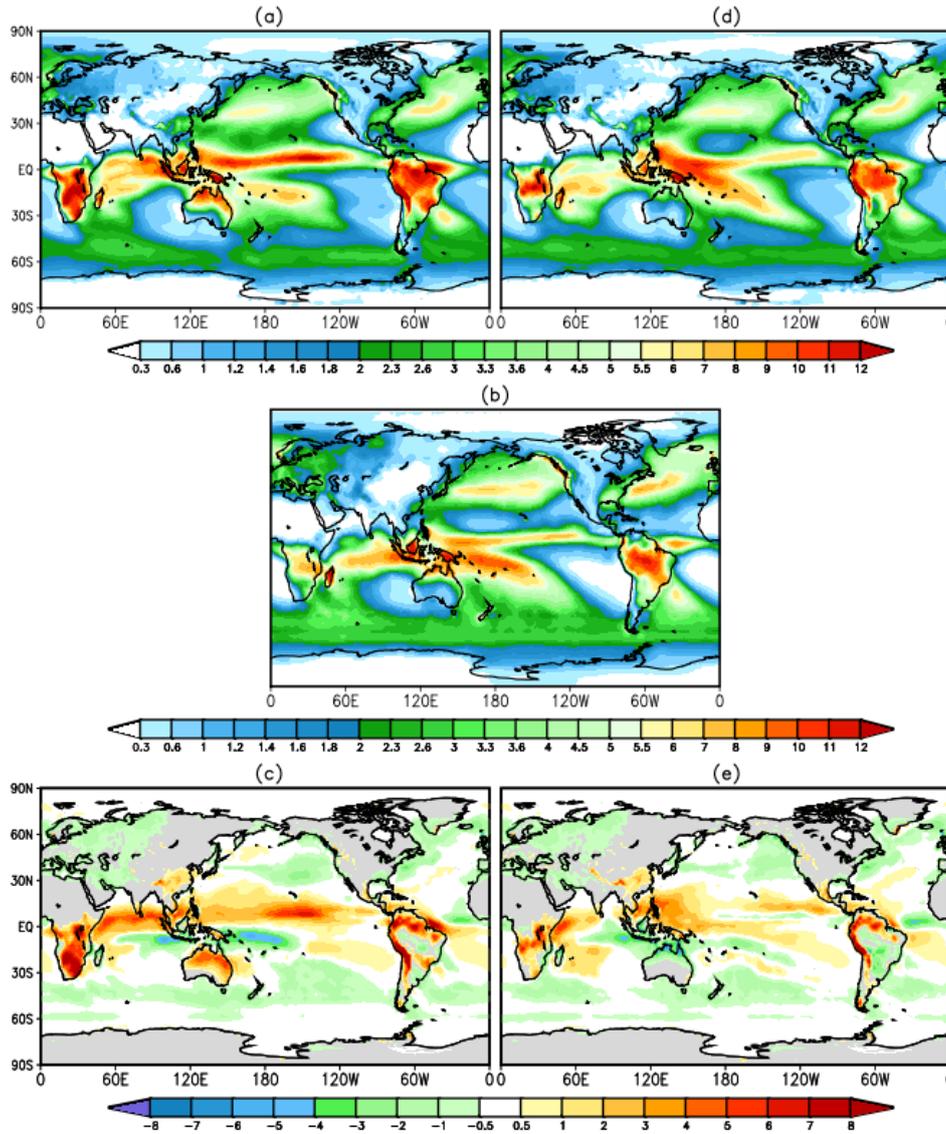


Figure 12. December-January-February total precipitation in mm day^{-1} from (a) experiment 6 (MERRA AGCM-like), (b) GPCP, (c) the difference of experiment 6 minus GPCP, (d) experiment 5 (MERRA2 AGCM-like), (e) the difference of experiment 5 minus GPCP. The gray shading indicates land areas.

change from the MERRA2 AGCM to the MERRA AGCM re-evaporation, and, as expected, shows the drying related to the reduced re-evaporation in the MERRA AGCM-like experiment. Figure 11a–c show the change in total precipitable water, and the difference (MERRA AGCM formulation minus MERRA2 AGCM formulation, shown in Fig. 11c) is always negative everywhere. The vertical distribution of the moisture is shown with the specific humidity in Fig. 11d–f, where the difference plot (Fig. 11f) also shows an almost global reduction in atmospheric water vapor.

The resulting mean circulation in boreal winter underwent a substantial change associated with this drying, and represents the most substantial impact on the simulated

climate of all the elements of the MERRA AGCM to MERRA2 AGCM transition. Figure 12 shows the substantial impact that changing the moisture levels had on the 30-year-averaged December-January-February total precipitation. Given the Global Precipitation Climatology Project (GPCP; Huffman et al., 1995) climatology as a reference, the MERRA2 AGCM-like experiment (experiment 5, Fig. 12d) exhibits an Intertropical Convergence Zone (ITCZ) structure that is properly placed in longitude, a South Pacific Convergence Zone (SPCZ) that properly slants from the western Pacific to the southeast, and more accurately simulated storm track maxima. The precipitation field represents improvements in all these areas relative to the MERRA AGCM-like

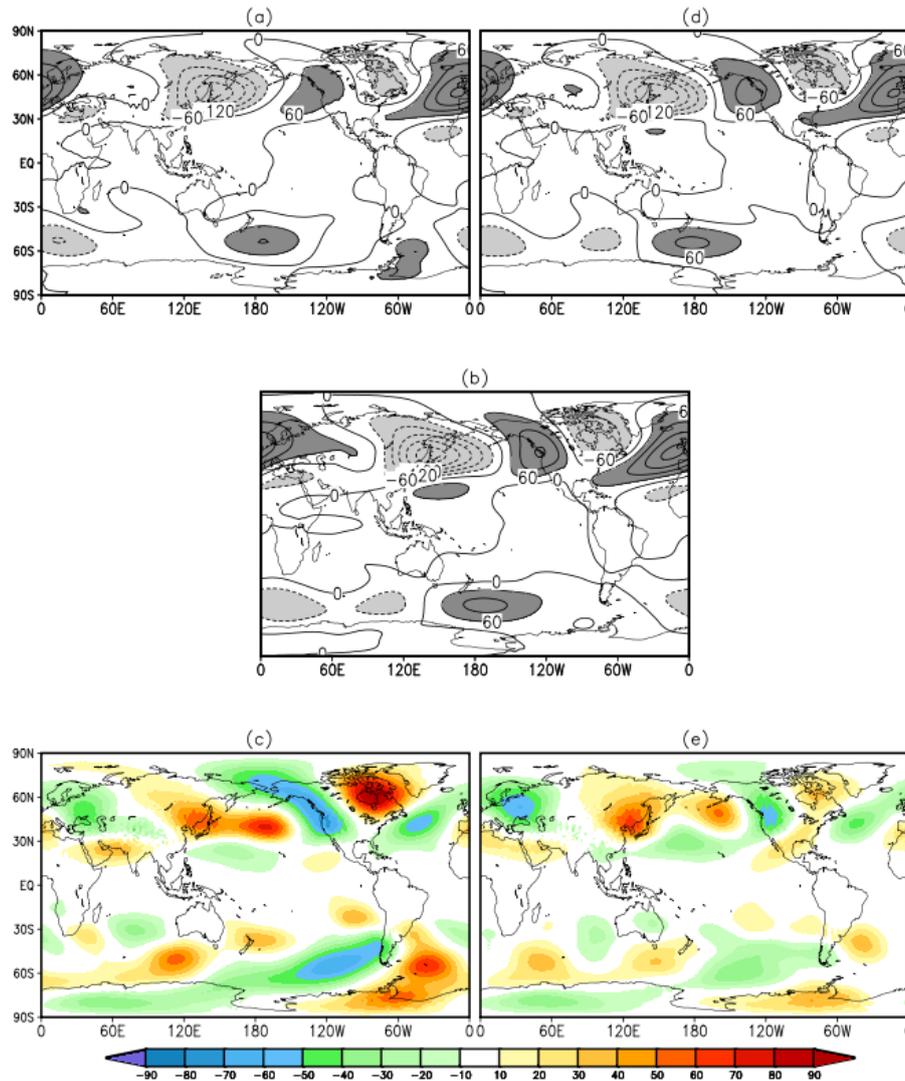


Figure 13. December-January-February 300 mb eddy height climatology in meters from (a) experiment 6 (MERRA AGCM-like), (b) MERRA, (c) the difference of experiment 6 minus MERRA, (d) experiment 5 (MERRA2 AGCM-like), (e) the difference of experiment 5 minus MERRA.

simulation (experiment 6) shown in Fig. 12a. The change in the tropical precipitation field related to the additional drying in experiment 6 also resulted in substantial changes in the Pacific teleconnection patterns, specifically the Pacific–North America (PNA) pattern in boreal winter. The eddy height field is an indicator of the strength of the PNA, and is shown in Fig. 13 in relation to the eddy height from MERRA reanalysis. The MERRA2 AGCM-like simulated PNA pattern (Fig. 12d) has a stronger and more properly oriented ridge near the west coast of North America relative to the PNA as simulated by the MERRA AGCM-like experiment (Fig. 12a). This change has implications for the poleward propagation of heat and momentum. The standard deviation of the difference from MERRA reanalysis estimates is also substantially reduced in experiment 5 (16.8 m) relative to ex-

periment 6 (20.3 m). The direct impact of the change in re- evaporation was also evident in the boreal summer climatology. Figure 14 shows this both in the total precipitable water fields and in the specific humidity fields, where the differences between experiments 5 and 6 are of the order of the differences seen in boreal winter. The impact on the mean summertime circulation, however, was minimal.

3.5 Breakup of the Southern Hemisphere stratospheric jet

The changes to the gravity wave drag parameterization included the modification of the intermittency factor, used to reduce the strength of the gravity wave drag based on expected departure from linear theory. The value of the intermittency factor was increased in the MERRA2 AGCM for

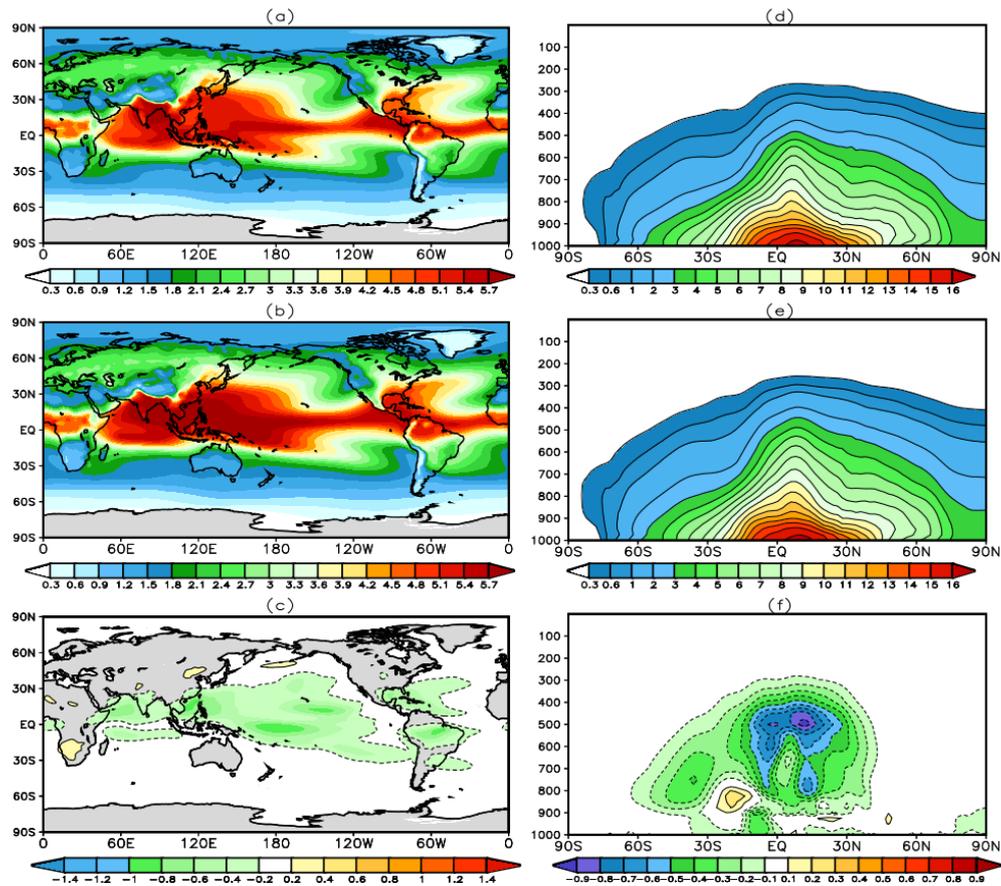


Figure 14. (a) June–July–August total precipitable water in mm from experiment 6 (MERRA AGCM-like), (b) same as (a) but from experiment 5, (c) same as (a) but the difference of experiment 6 minus experiment 5, (d) June–July–August-specific humidity in g kg^{-1} from experiment 6 (MERRA AGCM-like), (e) same as (d) but from experiment 5, (f) same as (d) but the difference of experiment 6 minus experiment 5.

orographic waves as a function of latitude, changing from a MERRA AGCM global value of 0.125 to values reaching 0.3125 south of approximately 40°S . This behavior is based on observational evidence of strong gravity waves from isolated small mountains in the Antarctic peninsula (Alexander and Teitelbaum, 2007) and the South Georgia islands (Alexander et al., 2009).

Experiment 7 is the last experiment in the series, incorporating the effects of all the fundamental changes between the MERRA and MERRA2 versions of the GEOS-5 AGCM. The removal of the intermittency change is expected to decrease the orographically induced drag in the Southern Hemisphere, thereby depositing less momentum aloft and increasing the strength of the westerlies. The zero-wind contour of the zonal wind averaged from 50 to 70°S in the Southern Hemisphere can be used as an indicator of the level and timing of the stratospheric jet breakup. Figure 15 shows the zero-wind contour of the averaged zonal wind from experiment 7, experiment 6, and MERRA. At the 2-degree horizontal resolution of the simulations described here, the zero-

wind line is higher in altitude and delayed in time relative to MERRA in both simulations, reflecting a delayed stratospheric jet breakup. The decreased intermittency factor in experiment 7, however, delays the jet breakup even more, showing the improvement in the MERRA2 AGCM-like simulation relative to the MERRA AGCM-like simulation. At higher spatial resolution (not shown) the increase of intermittency factor in the MERRA2 AGCM is effective in producing a reasonable evolution of the polar vortex breakdown in the Southern Hemisphere.

4 High-resolution simulations and resolution-aware behavior in the MERRA2 AGCM

The modifications to the MERRA2 AGCM physical parameterizations described above resulted in improvements in simulated climate at all the resolutions relevant to GMAO. Additional developments were implemented in the MERRA2 AGCM that were particularly applicable to higher-resolution (0.25° or higher) simulations. These included the changes in

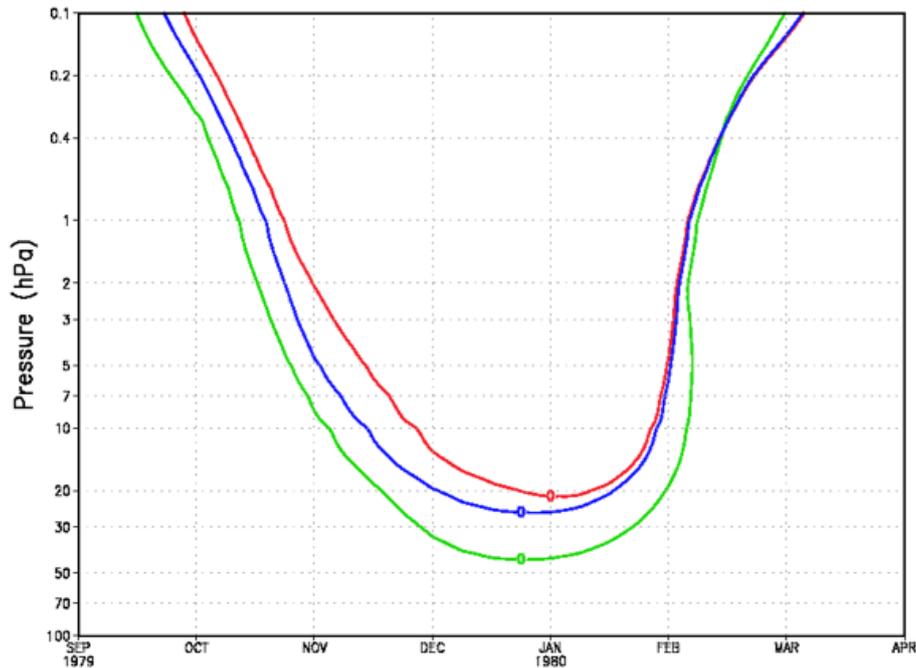


Figure 15. The annual cycle of the zero contour of the 30-year-averaged zonal mean zonal wind, averaged from 70 to 50° S latitude in m s^{-2} from experiment 7 in red (MERRA AGCM-like), experiment 6 in blue (MERRA2 AGCM-like) and MERRA in green.

ocean surface roughness at high wind speeds (mentioned in Sect. 3.1 and examined in detail in Molod et al., 2013) and the implementation of resolution-aware parameters.

The implementation of the resolution-aware behavior of the moist processes in the MERRA2 version of the GEOS-5 AGCM was designed in part to improve the behavior of the high-resolution simulations and to ensure more uniformity of model mean state across resolutions and applications. Figure 16 shows an example of the undesirable behavior in the MERRA version of the AGCM that the resolution-aware parameters were meant to address. Figure 16a–c shows the difference between MERRA AGCM and ERA-Interim specific humidity at 2., 1. and 0.5° resolution, indicating the increase of the error in atmospheric moisture content relative to ERA-Interim reanalysis (Dee et al., 2011) as the resolution increases.

Studies with many AGCMs used for climate and weather applications have reported large changes in the model solution as the horizontal resolution increases. For example, Williamson (2008) showed sensitivity of aqua-planet simulations to both resolution and time step, Roecker et al. (2006) reported on lack of convergence toward a single solution as the horizontal and vertical resolution changed even though the higher-resolution simulations showed improvements relative to the coarse resolution result, and Bacmeister et al. (2014) showed a spurious precipitation pattern when the AGCM resolution was increased to 0.25°.

In order to mitigate the change (increasing error) in the GEOS-5 AGCM simulated climate with increasing resolu-

tion, the MERRA2 AGCM moist physics was modified to include two governing parameters that are specified a priori as a function of horizontal resolution. These are the RH_{crit} used for large-scale condensation, and a parameter which governs the minimum allowable entrainment used for the stochastic Tokioka trigger of the convective parameterization.

The dependence of RH_{crit} on horizontal resolution is based on the Molod (2012) analysis of a global mesoscale simulation, and is such that the RH_{crit} above the planetary boundary layer (PBL) increases with finer resolution, as is seen in the progression from the 2° resolution curve up to the 0.125° curve in Fig. 7. This progression is consistent with an intuitive expectation that the variability of total water within an AGCM grid cell decreases as the grid cell becomes smaller. The implementation of the horizontal resolution dependence of RH_{crit} (as described in Sect. 2) in the MERRA2 AGCM resulted in an atmospheric moisture field that is more consistent across different resolutions, as seen in Fig. 16d–f.

The MERRA2 version of the GEOS-5 AGCM also includes a horizontal resolution-dependent and stochastic Tokioka-type trigger (described in Bacmeister and Stephens, 2011) as part of the RAS convective parameterization. The cloud model in RAS computes the effect of individual entraining cloud plumes, and the trigger acts to effectively eliminate any cloud plume with too small an entrainment during ascent. Bacmeister and Stephens (2011) examined the observed relationship between the neutral buoyancy level of a particular sounding and the observed condensate (a proxy for the convective detrainment level), and found that this ob-

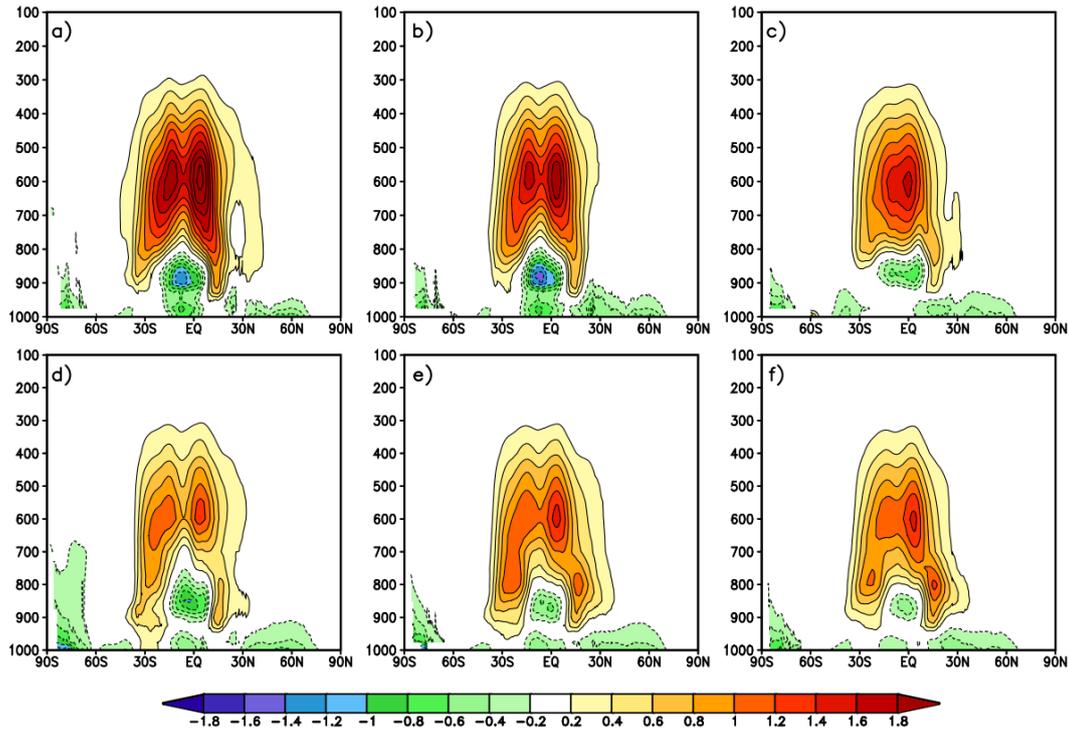


Figure 16. December-January-February seasonal mean difference of specific humidity in g kg^{-1} for (a) MERRA-like AGCM – ERA-Interim at 2-degree resolution, (b) MERRA-like AGCM – ERA-Interim at 1-degree resolution, (c) MERRA-like AGCM – ERA-Interim at 1/2-degree resolution, (d) MERRA2-like AGCM – ERA-Interim at 2-degree resolution, (e) MERRA2-like AGCM – ERA-Interim at 1-degree resolution, (f) MERRA2-like AGCM – ERA-Interim at 1/2-degree resolution.

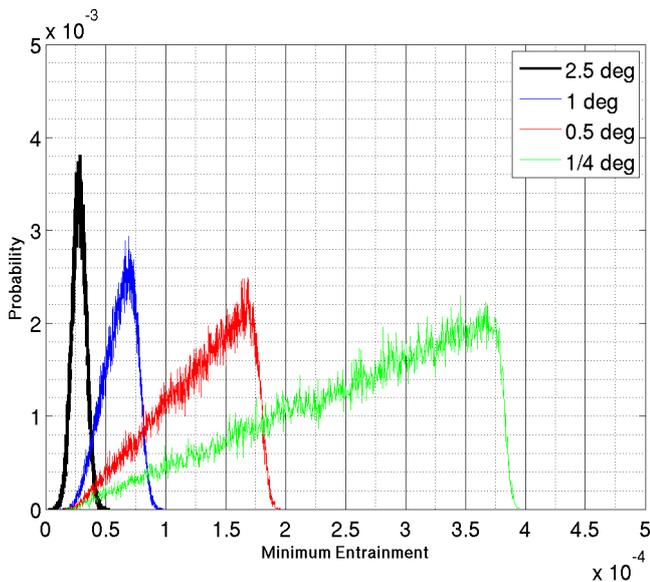


Figure 17. The PDF for the minimum entrainment allowed by the cumulus parameterization for different AGCM horizontal resolutions. Black line is for 2×2.5 degrees, blue for 1 degree, red for 1/2 degree and green for 1/4 degree.

servational relationship can be approximated by sampling the minimum entrainment from a power law PDF. The parameters of the PDF are specified a priori in the MERRA2 AGCM, and vary with resolution. The stochasticity is designed to only occasionally permit the least entraining and therefore the deepest detraining cloud plumes. The resolution dependence of the stochastic Tokioka trigger is such that it more severely limits the parameterized convective mass flux at high spatial resolution, where we expect the larger scale convective updrafts to be resolved explicitly, and has little impact at low resolution. This approach enables a relatively smooth transition from a fully parameterized cumulus convection at coarse resolution to a fully resolved cumulus convection at extremely high resolution. The ad hoc specified change with resolution of the PDF parameters is shown in Fig. 17.

The basic approach implemented in the MERRA2 AGCM to limit the impact of the parameterized convection at higher resolution is consistent with the approach of Arakawa and Wu (2013). They remove the assumption that the fraction of the grid box covered by convection is small, and model the total convective transport in terms of a parameterized contribution and a contribution from eddy transport. The relative importance of these two contributions changes with resolution and in this way they generalize the cumulus parameteri-

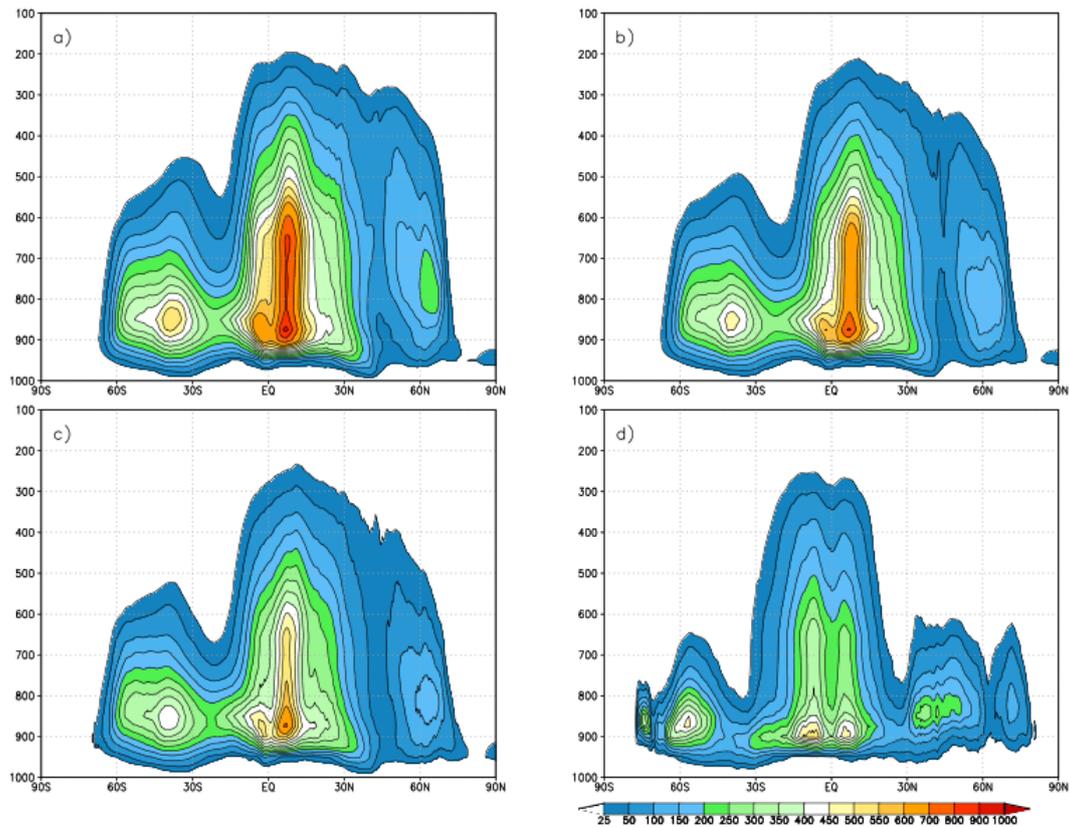


Figure 18. June-July-August seasonal mean cumulus mass flux in $\text{kg m}^{-2}\text{day}^{-1}$ from MERRA2-like AGCM for (a) 1-degree resolution, (b) 1/2-degree resolution, (c) 1/4-degree resolution, (d) 10 km resolution.

zation so that it converges to an explicit simulation of convection at high resolution. Another approach to implementing resolution-aware behavior in an AGCM cumulus parameterization is that of Bechtold et al. (2008), implemented in the Integrated Forecast System Model Cycle 40r1 of the European Center for Medium Range Weather Forecasting (documentation available at <http://old.ecmwf.int/research/ifsdocs/>). They compute a cumulus timescale that is sampled from a PDF whose parameters change with resolution, such that at high resolution the probability of finding a short timescale is increased. The net result is that the relative contribution of convective and resolved precipitation is relatively unchanged across resolutions, but this scheme does not converge to allow a total dependence on explicitly resolved convection at the highest resolutions.

Figure 18 shows a sequence of the June-averaged convective mass fluxes from MERRA2 AGCM simulations with different horizontal resolutions. The decrease of parameterized mass flux with resolution reflects the increasingly restrictive trigger, selected from the PDFs with increasingly higher minimum entrainment values shown in Fig. 17. The effect of this repressed RAS mass flux on the simulated climate is reflected in the total change in moisture due to moist processes, shown in Fig. 19. The cloud model in RAS includes a grid-scale sub-

sidence to compensate for the updraft mass flux that results in a drying of the sub-cloud layer. Figure 19 shows the decrease of the low level drying with increased horizontal resolution. The benefits of the reduced low level drying and the related maintenance of the cumulus available potential energy during tropical storm development in the GEOS-5 AGCM simulations at high resolution was demonstrated in the study of Lim et al. (2014).

5 Synthesis and discussion

The version of the GEOS-5 AGCM used in the Goddard Modeling and Assimilation Office (GMAO) MERRA2 re-analysis was developed for use across many different resolutions and applications. A unique series of AGCM simulations were performed with the GEOS-5 AGCM to detail the impact of each change in parameterization between the MERRA version and the current MERRA2 version. The series of sensitivity experiments began with the current AGCM version and regressed, one step of development at a time, to the MERRA AGCM.

The most substantial positive impact on the simulated AGCM climate was shown to be attributable to the increase of the re-evaporation of frozen cloud water and precipita-

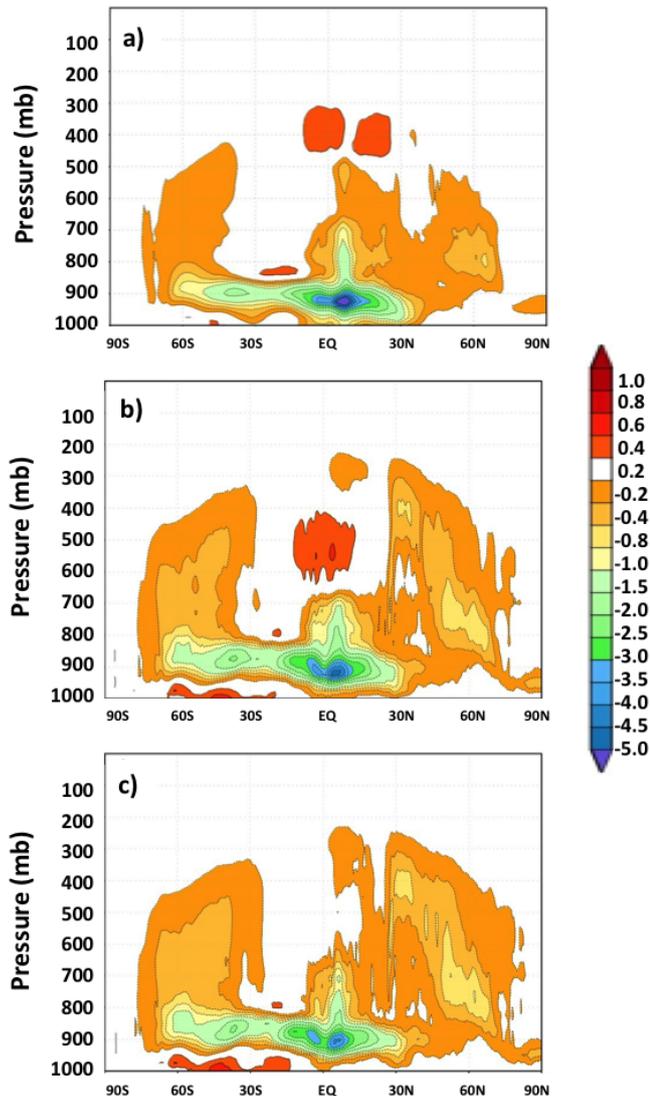


Figure 19. June–July–August seasonal mean change in moisture due to moist processes in $\text{kg m}^{-2} \text{day}^{-1}$ from MERRA2-like AGCM for (a) 1-degree resolution, (b) 1/2-degree resolution, (c) 1/4-degree resolution.

tion in the MERRA2 AGCM. The resulting atmosphere had a higher moisture content, and many aspects of the boreal winter climate were substantially improved relative to reanalysis. The moisture and cloud cover amounts were shown to be further improved by the implementation of an AIRS-based PDF of total water.

The development of the MERRA2 AGCM also included the implementation of a set of parameters governing moist processes that contain an a priori change in behavior with horizontal resolution. The parameters are ones which govern the minimum allowable entrainment into a convective up-draft, and govern the atmospheric RH needed for the onset of condensation. Results of a limited set of experiments were

shown to demonstrate the benefits of this resolution-aware behavior at higher resolution.

This study was focused on the results of atmosphere-only simulations, but the resulting model has also performed well in coupled atmosphere–ocean, coupled atmosphere–chemistry, data assimilation, numerical weather prediction, and global mesoscale applications.

Code availability

The GEOS-5 source code is available under the NASA Open-Source Agreement at <http://opensource.gsfc.nasa.gov/projects/GEOS-5/>.

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