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Supplement of

A soil diffusion–reaction model for surface COS flux: COSSM v1

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To demonstrate quantitatively that aqueous diffusive flux is small compared to gaseous diffusive flux, we assume a soil column of vertically uniform profiles of porosity, temperature and moisture,

$$\theta_{\text{sat}} = 0.50 \text{ m}^3 \text{ m}^{-3}$$

$$T = 25^\circ\text{C}$$

$$\theta_{\text{w}} = 0.25 \text{ m}^3 \text{ m}^{-3}$$

The factor b in Eq. (13) in the main text is set to 5.3 here, without losing generality. The soil gaseous diffusivity for COS calculated based on Eq. (13) in the main text is, $D_{\text{g}} = 5.64 \times 10^{-7} \text{ m}^2 \text{ s}^{-1}$. The molecular aqueous diffusivity for COS calculated from Ulshöfer et al. (1996) is $D_{\text{m, aq}} = 1.94 \times 10^{-9} \text{ m}^2 \text{ s}^{-1}$. Then by counting the tortuosity effect using Eq. (3) in Moldrup et al. (2003), we obtain the actual soil aqueous diffusivity for COS, $D_{\text{aq}} = 7.13 \times 10^{-11} \text{ m}^2 \text{ s}^{-1}$.

We consider the steady-state COS flux of the soil (i.e. $\partial C/\partial t=0$) with a sink proportional to the concentration and no production term. We obtain the equation,

$$D \frac{d^2 C}{dz^2} = -S = kC \quad (1)$$

with the boundary condition $C(z=0) = C_{\text{atm}}$. We obtain the general solution,

$$C(z) = C_{\text{atm}} \exp(-\sqrt{k/D} \cdot z) \quad (2)$$

The surface flux is thus

$$F = -D \left. \frac{dC}{dz} \right|_{z=0} = \sqrt{kD} C_{\text{atm}} \quad (3)$$

which is proportional to \sqrt{D} . We can then calculate the ratio of aqueous diffusive flux (F_{aq}) to gaseous diffusive flux (F_{g}),

$$\frac{F_{\text{aq}}}{F_{\text{g}}} = \sqrt{\frac{D_{\text{aq}}}{D_{\text{g}}}} = 0.0112 \quad (4)$$

Under most conditions, the aqueous diffusive flux is a small fraction of the total flux (Fig. S1). Hence, any errors from neglecting aqueous diffusion are usually small (< 10% at water filled pore space below 83%), except at high soil moisture when the overall flux also tends to be small (see Fig. 9).

References

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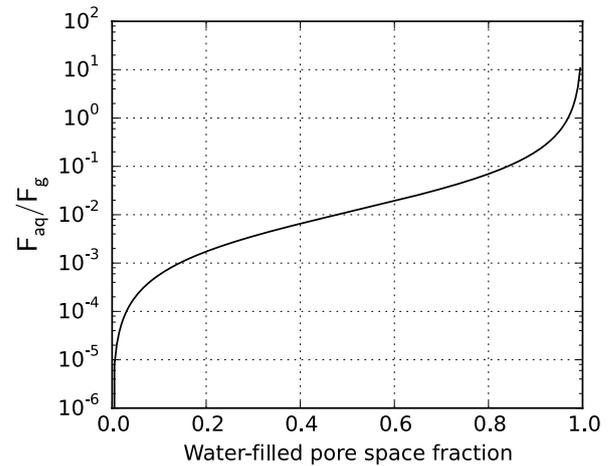


Figure S1. The ratio of aqueous flux to gaseous flux for an ideal case with soil porosity $0.50 \text{ m}^3 \text{ m}^{-3}$. COS uptake velocity is assumed constant.